

**THE TEMPORAL STABILITY AND ACTIVITY  
OF LANDSLIDES IN EUROPE  
WITH RESPECT TO CLIMATIC CHANGE**

CEC Environment Programme - Contract N° EV5V-CT94-0454

**1st TESLEC MEETING  
Padua, Italy**

10th - 12th July 1994

Local organizers:

**National Research Council  
Research Institute for the Prevention  
of Geological and Hydrological Hazard  
Padua**

## SUMMARY OF THE PROGRAMME

### SUNDAY 10th July

#### Opening Session

14.00 - 15.30	Structure and organization of the project and of the meeting (Budget, deliverables, deadlines, responsibilities, addresses, distribution of the handbook etc.)
15.30 - 16.00	Coffee break

#### Afternoon Session

#### Objective 1: RECOGNITION OF RELICT LANDSLIDES

16.00 - 19.00	The Manual - essential tasks (Key-note: Denys Brunsdén) - <i>Methods (remote sensing, mapping, classification, description of the behaviour)</i> - <i>Checklist for non-specialists</i>  Deadlines and responsibilities
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Evening free

### MONDAY 11th July

#### Morning Session I

Objective 2:	<b>PAST DISTRIBUTIONS OF LANDSLIDE INCIDENTS AND THEIR RELATIONSHIP TO CLIMATIC CHANGE PARAMETERS. QUALITATIVE EVOLUTION MODELS (LANDSLIDE AND LANDSCAPE EVOLUTION)</b>
09.00 - 10.30	Discussion of national records of landslides and times series - <i>Data techniques, central database of the project (data availability and exchange)</i>
10.30 - 11.00	Coffee break

#### Morning Session II

11.00 - 12.30	Discussion of test sites - <i>Reports on the proposed National Programmes (max. 15-20 min: France/Spain/Italy)</i> - <i>Central database (data availability and exchange)</i>  Deadlines and responsibilities
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#### Objective 3: HYDROLOGICAL AND SLOPE STABILITY MODELLING

12.30 - 12.45	Introduction (Richard Dikau)
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12.45 - 13.15 Hydrological models (Key-note: Theo van Asch, max. 30 min)  
- *Discussion*

13.15 - 15.00 Lunch at C.N.R. Canteen

### Afternoon Session I

15.00 - 15.45 Slope stability models (Key-note: Eddie Bromhead, max. 30 min)  
- *Discussion*

15.45 - 16.30 Climate change models (Key-note: Hans von Storch, max. 30 min)  
- *Discussion*

16.30 - 17.00 Coffee Break

### Afternoon Session II

17.00 - 19.00 General modelling framework  
- *Model inventory and link between hydrological, slope stability and climatic change models*  
- *Model application, development and exchange*

Deliverables and responsibilities

20.30 Meeting Dinner

## TUESDAY 12th July

### Morning Session

09.00 - 11.00 Workshops on modelling framework  
- *Constitution of working groups on hydrological models, slope stability and climatic change models*  
- *Discussion of scientific and technical problems*

11.00 - 11.30 Coffee Break

### Closing Session

11.30 - 13.00 Final discussion  
- *Summing up of the results and agreements (manual, data records, modelling framework)*  
- *Next workshops and meeting*

13.15 Lunch at C.N.R. Canteen

14.00 Departure

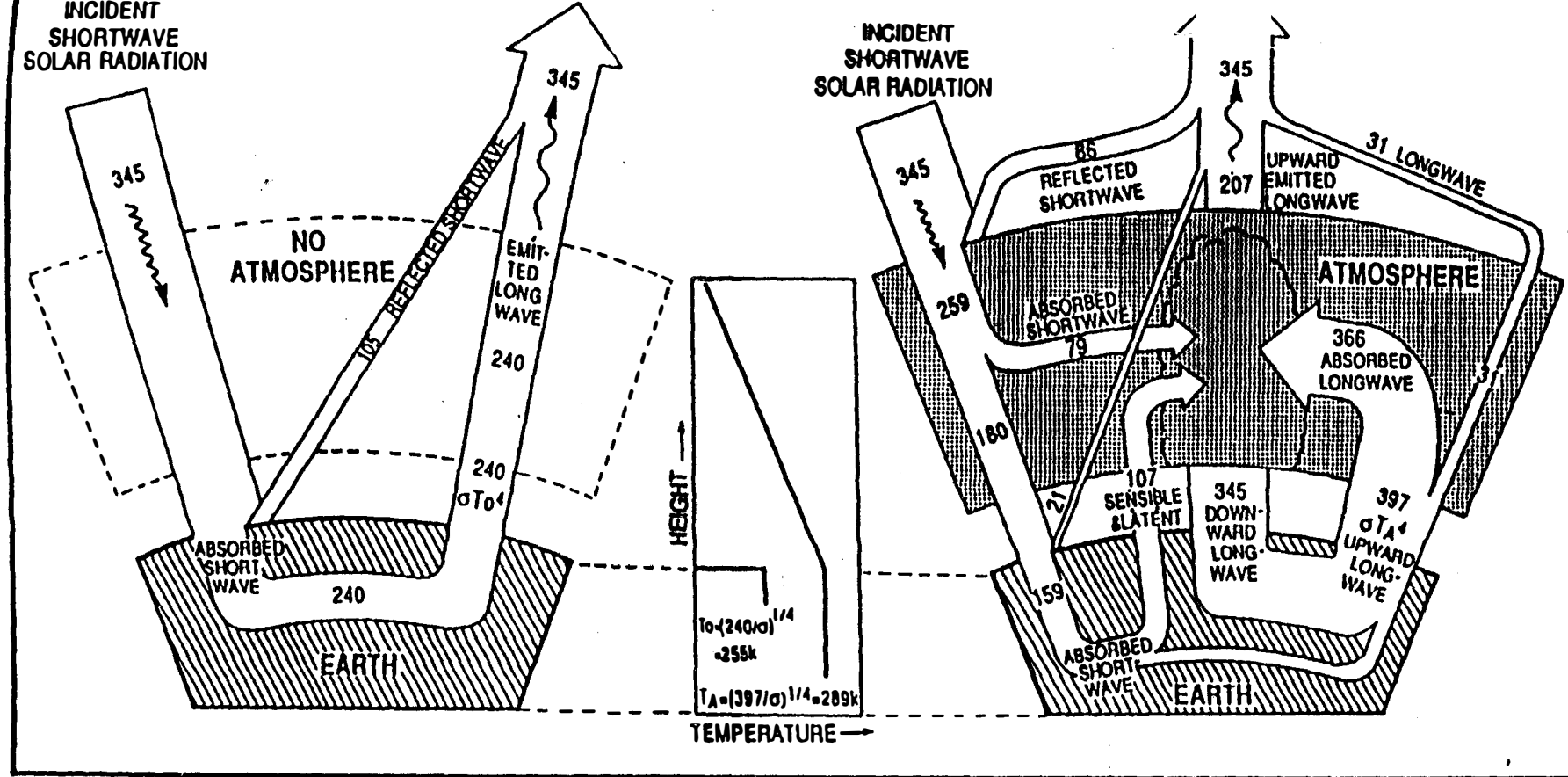
# Climate Change Models

**Hans von Storch**  
**Max-Planck-Institut für Meteorologie**  
**Hamburg, Germany**

Talks presented at the 1st TESLEC Meeting  
in Padua, Italy, on July 11 - 12, 1994

# Prologue

- **The Max-Planck-Institut für Meteorologie** is one of the centers where scenarios of global climate change are derived by means of coupled ocean/atmosphere/sea-ice dynamical models in
  - equilibrium runs (e.g, “ $2 \times CO_2$ ”)
  - “transient” runs (e.g., Scenario A)
- **These scenarios are used as input for climate impact studies** - for a responsible use in such studies the communication between scenario producers and scenario users is required.

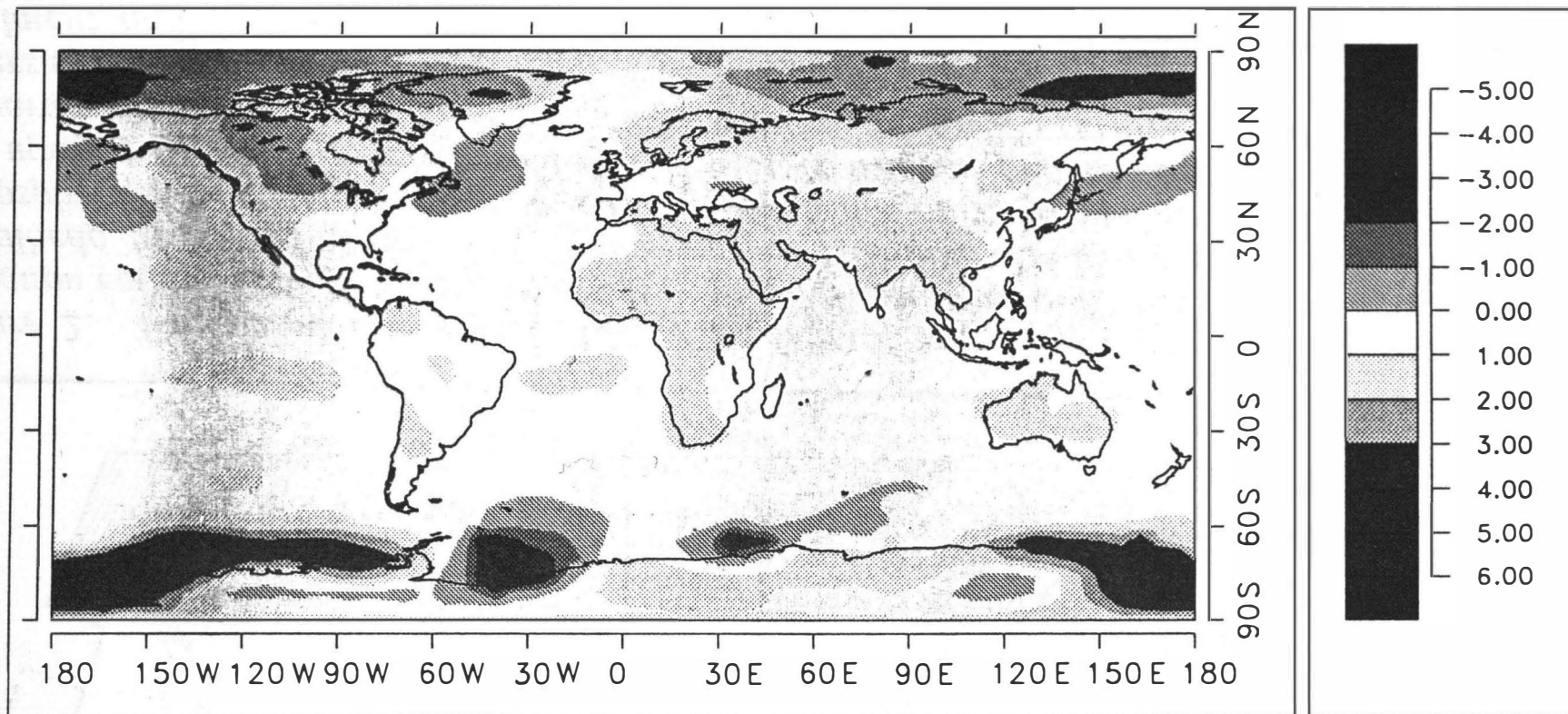


$$R_L \sim T^4$$

Figure 2. The greenhouse effect. In the absence of an atmosphere (left), the long-wave radiation emitted back to space will exactly equal the absorbed short-wave radiation (assuming an albedo of approximately 0.31) if the surface radiative temperature is 255K. With an atmosphere containing water vapour, carbon dioxide and other greenhouse gases (right) most of the upward radiation emitted from the surface is trapped and re-radiated downward and upward by cloud and greenhouse gases thus warming the earth's surface and cooling the upper layers of the atmosphere. The net long-wave radiative cooling of the atmospheric layer is offset by latent and sensible heat flux upward from the warmed surface. Radiative-convective equilibrium is achieved when the globally averaged surface temperature reaches 289K. In the equilibrium situation, the atmospheric temperature falls off with height to temperatures well below the planetary temperature of 255K as shown schematically above (centre). All heat fluxes are approximate global averages for the present climate and are given in Watts per square metre ( $\text{Wm}^{-2}$ )

(2026 - 2035) - (1986 - 1995)

2M TEMPERATURE CODE: 167 PRO: 91-100 EXPT: EIN



LABEL: (167) REF DATA: EIN YRS 51-60

GCH. MPI HH.

DATE: 04/17/94

TIME: 18:21:52

JMO-GRAPHICS

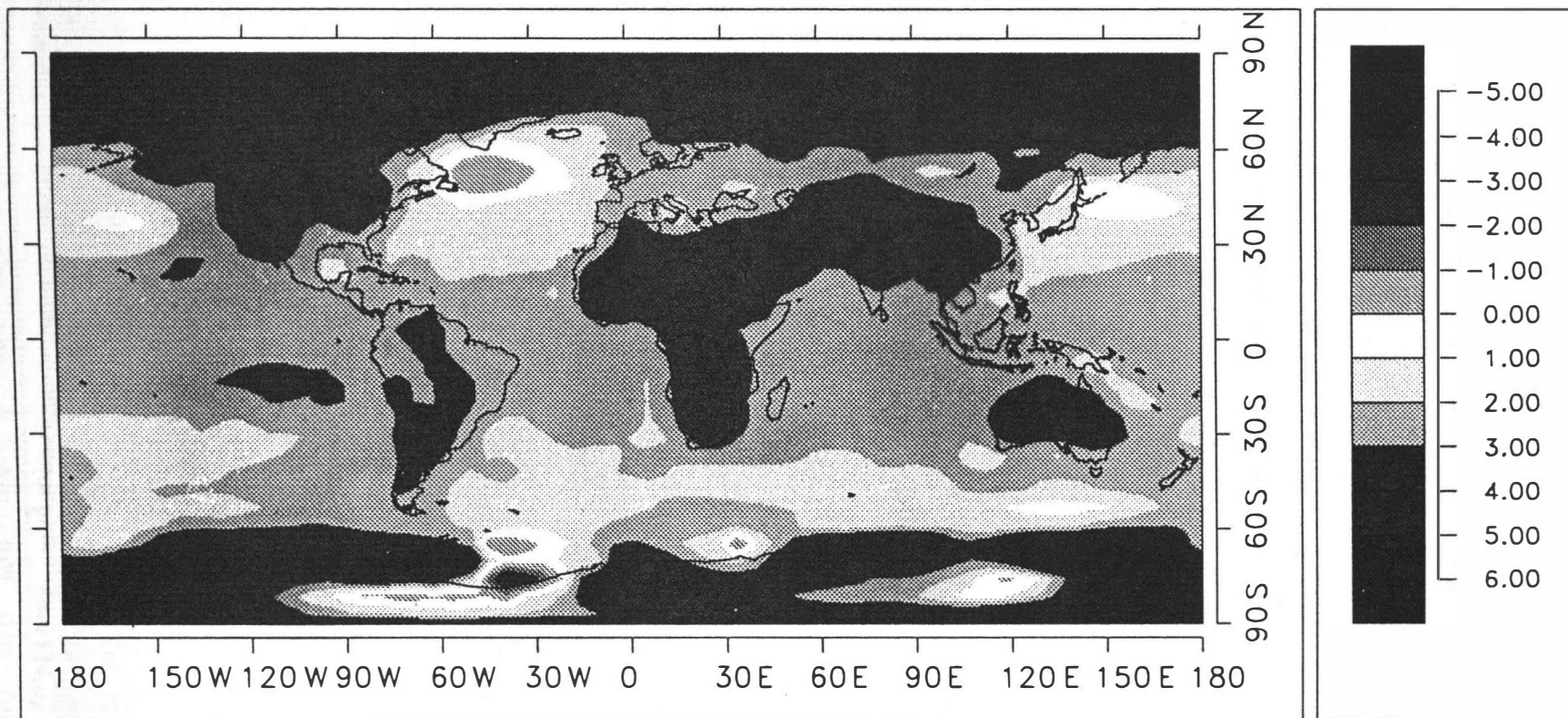
Von Cubarch et al, 1994

PRINT THIS SIDE

A4

(2076-85) - (1986-95)

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LABEL: (167) REF DATA: SZP YRS 51-60

GCH. MPI HH. DATE: 04/17/94

TIME: 18:21:17

JMO-GRAPHICS

*Von Cebasek et al, 1994*

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# Climate Change in the Past 100 Years - natural\* or external?

- The homogenized analyses of near-surface temperature indicate an increase of this temperature in the past 100 years.
- To assess whether this increase is natural or not a “detection variable” is formed, which is the projection of the analysed (“observed”) temperature fields on the warming pattern simulated by a climate model.
- The observed warming is declared “inconsistent with natural variability” (= signal has an external source) if the detection variable is outside the 95% confidence band of natural variability
- The 95% confidence band of natural variability is estimated from observed data (problem: contamination by the signal) or from climate model runs (problem: model might be wrong)
- This procedure leads to the decision

the recent warming trend is NOT natural.\*

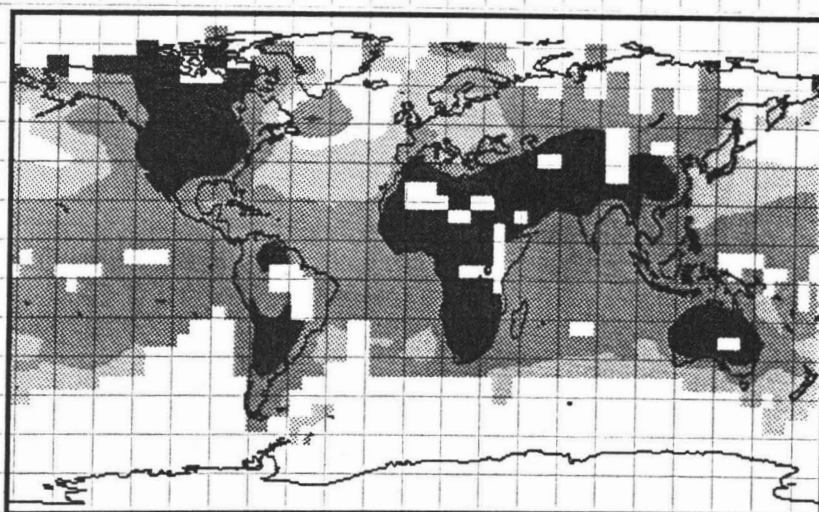
- An a-priori optimization of the “signal-to-noise” ratio reduces the risk of an incorrect assessment.

A publication is presently prepared by Hegerl et al. at the MPI in Hamburg

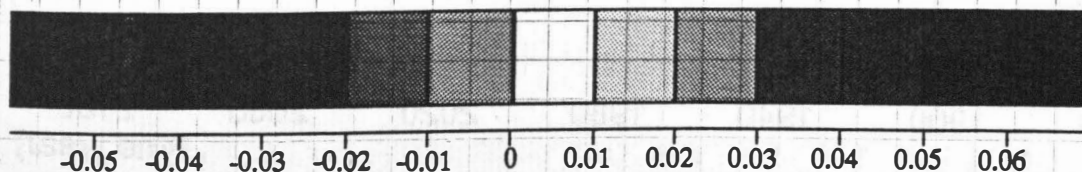
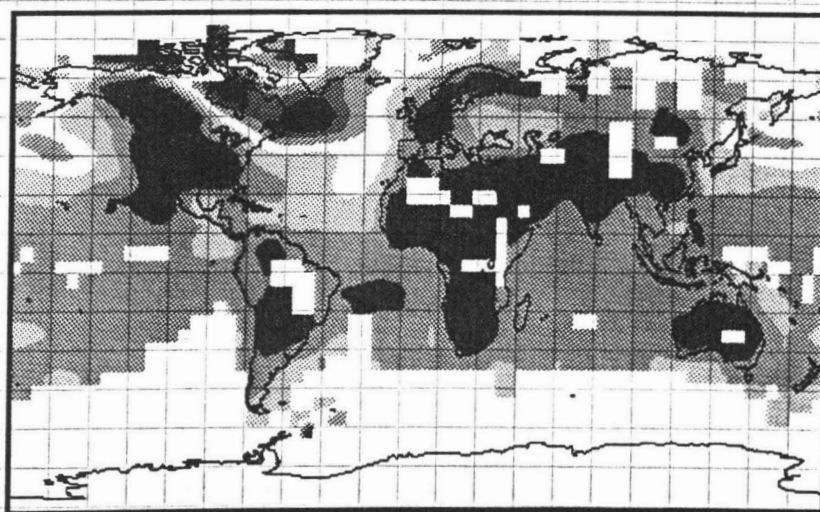
\*perhaps, “due to internal dynamics” would be a better phrase.

**Optimal Rotation for full set of Variability data**  
optimized in the space spanned by EOFs 1-4 of the  
Early Industrial global warming simulation

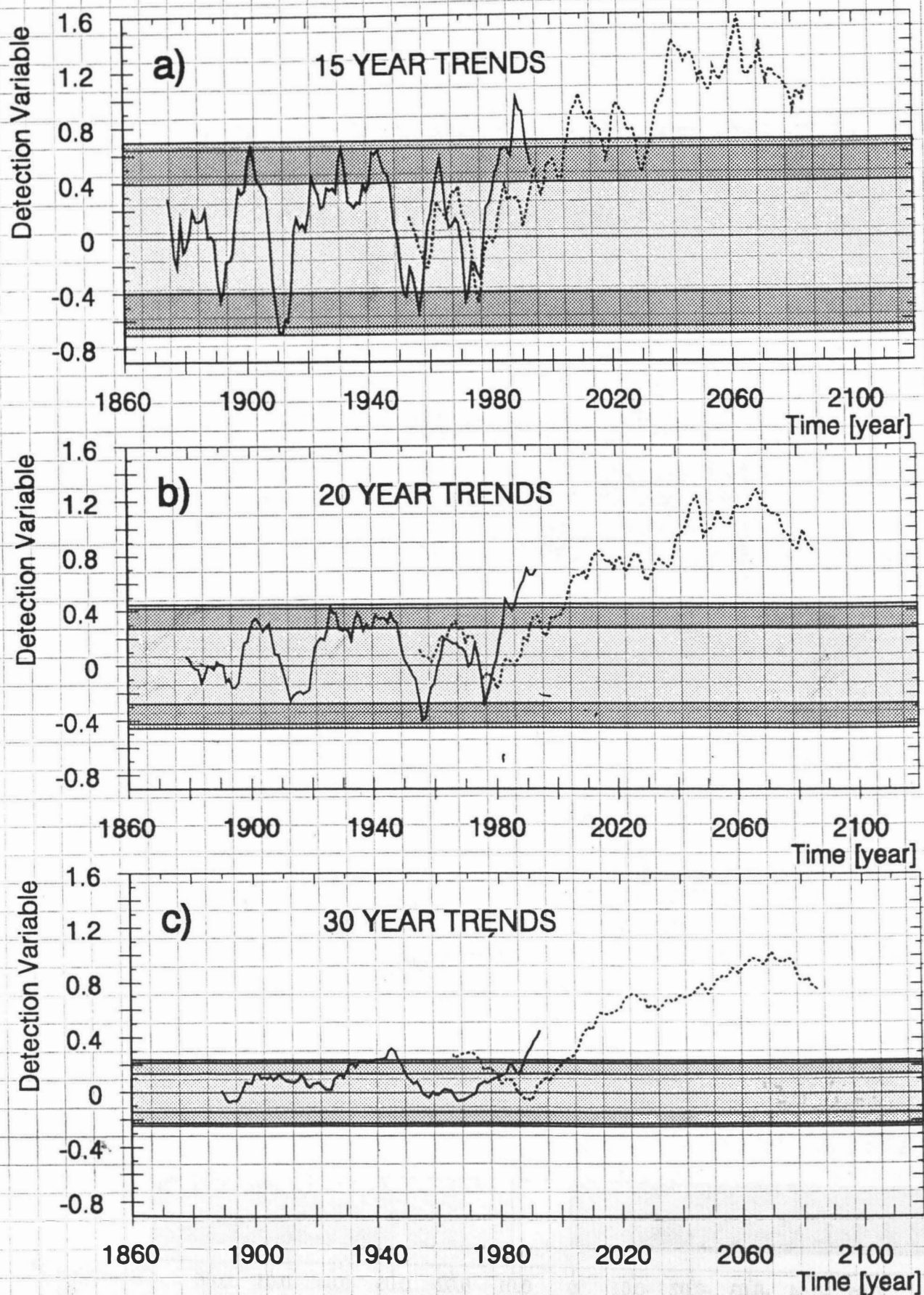
**guess-pattern  
fingerprint**



**optimally rotated  
fingerprint**



## SIGNIFICANCE OF THE OBSERVED PATTERN for the optimal fingerprint



- observed
- 95% confidence interval for the ECHAM / LSG control simulation
- 95% confidence interval for the ECHAM2 / OPYC control simulation
- 95% confidence interval of observations (GHG-signal subtracted)

# Climate Impact Research is NOT ...

- **The processing of static scenarios** in the sense that certain maps are given by the “producers” to the “users” without ongoing interaction. “Users” must know the conditions under which the scenarios are valid and the inherent limitation of such model generated data.
- **“CO<sub>2</sub>” scenarios.** The expected climate change is a gradual process. It will not be a sudden jump from one stationary state to another stationary state. That is, the important aspect is the rate of change.
- **The limitation to anthropogenic change.** Climate varies on all time and space scales because of internal dynamics. Also such internally excited changes affect climate-sensitive systems. Anticipated impacts due to man-made climate change should be compared to experienced impacts of past natural variations.

# GRAIN MAIZE

Suitability  
Change from  
1990-2050

Selections  
View map  
Zoom in  
Parameters  
Exit

Yes  
Extension  
Retreat  
No

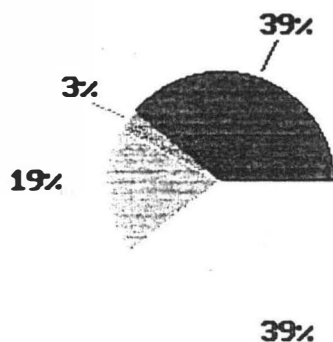


Policy:GLOBAL  
Ref :1990

STAGGER : best  
CLIMAPS : best

dT=1.26°C  
Year 2050

EC

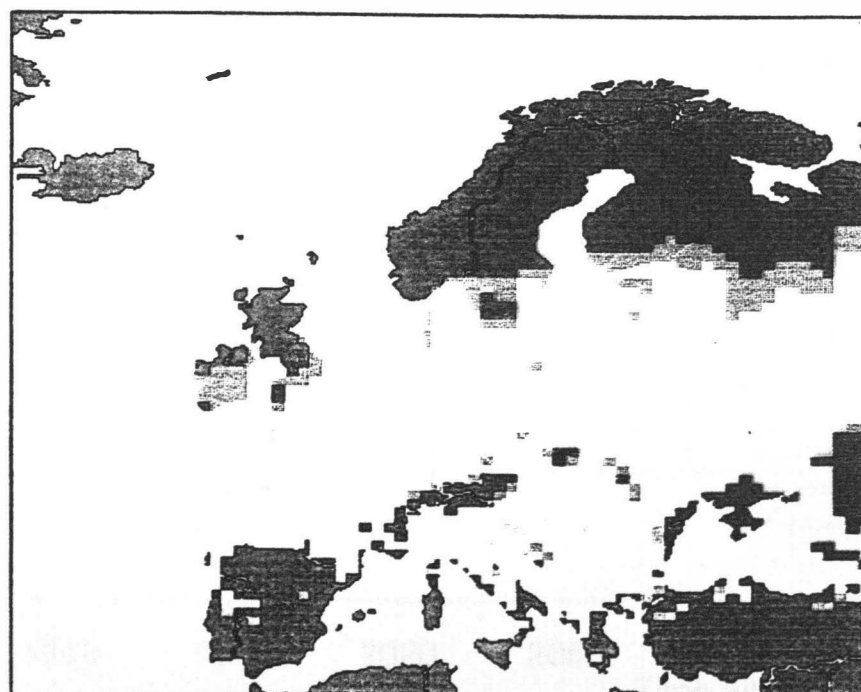


# GRAIN MAIZE

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Change from  
Ref to Pol.

Selections  
View map  
Zoom in  
Parameters  
Exit

Yes  
Extension  
Retreat  
No

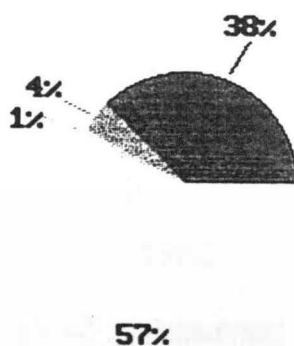


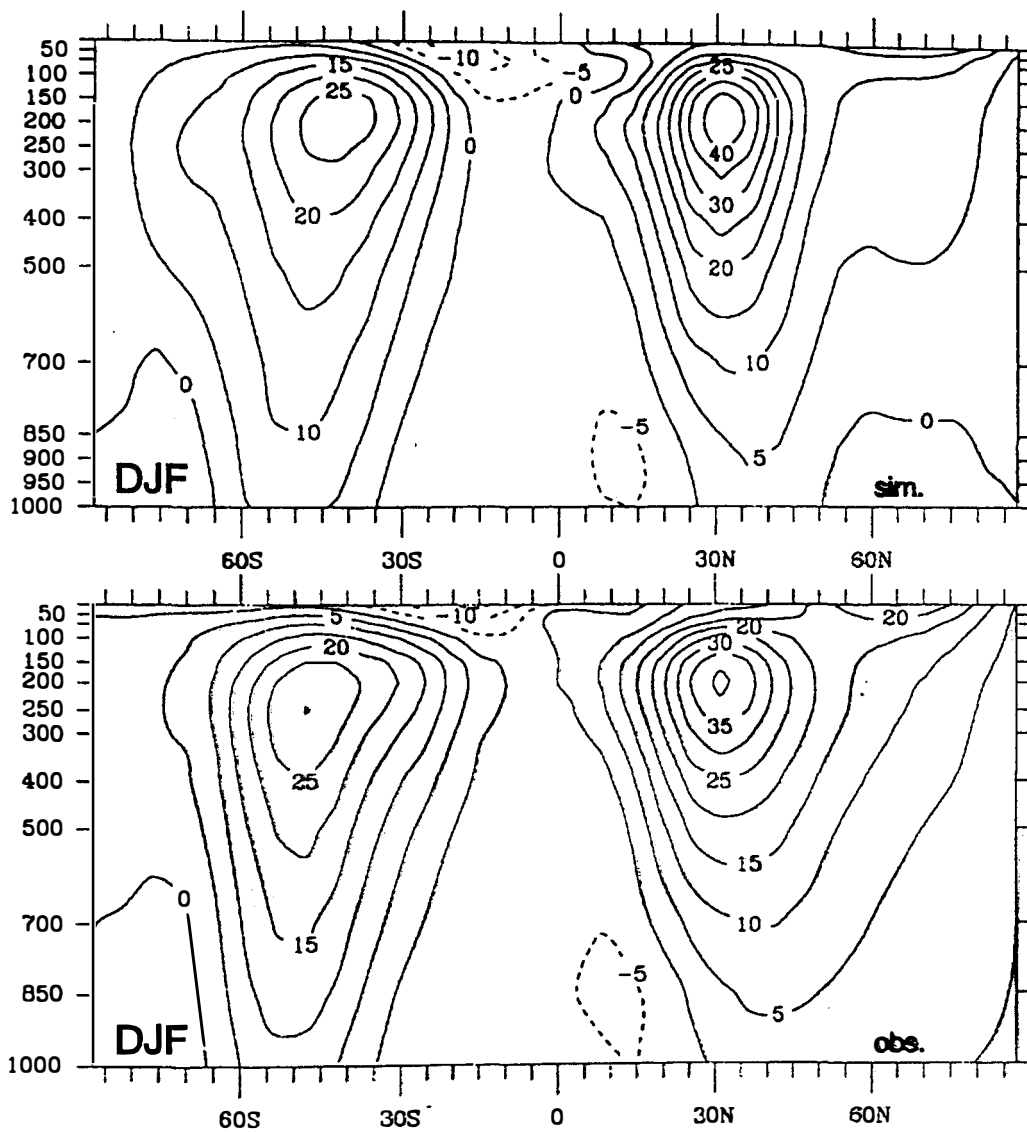
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Ref :RS1991A

STAGGER : best  
CLIMAPS : best

dT=-0.62°C  
Year 2050

EC





044HVSb.drw



Hulme et al., 1992

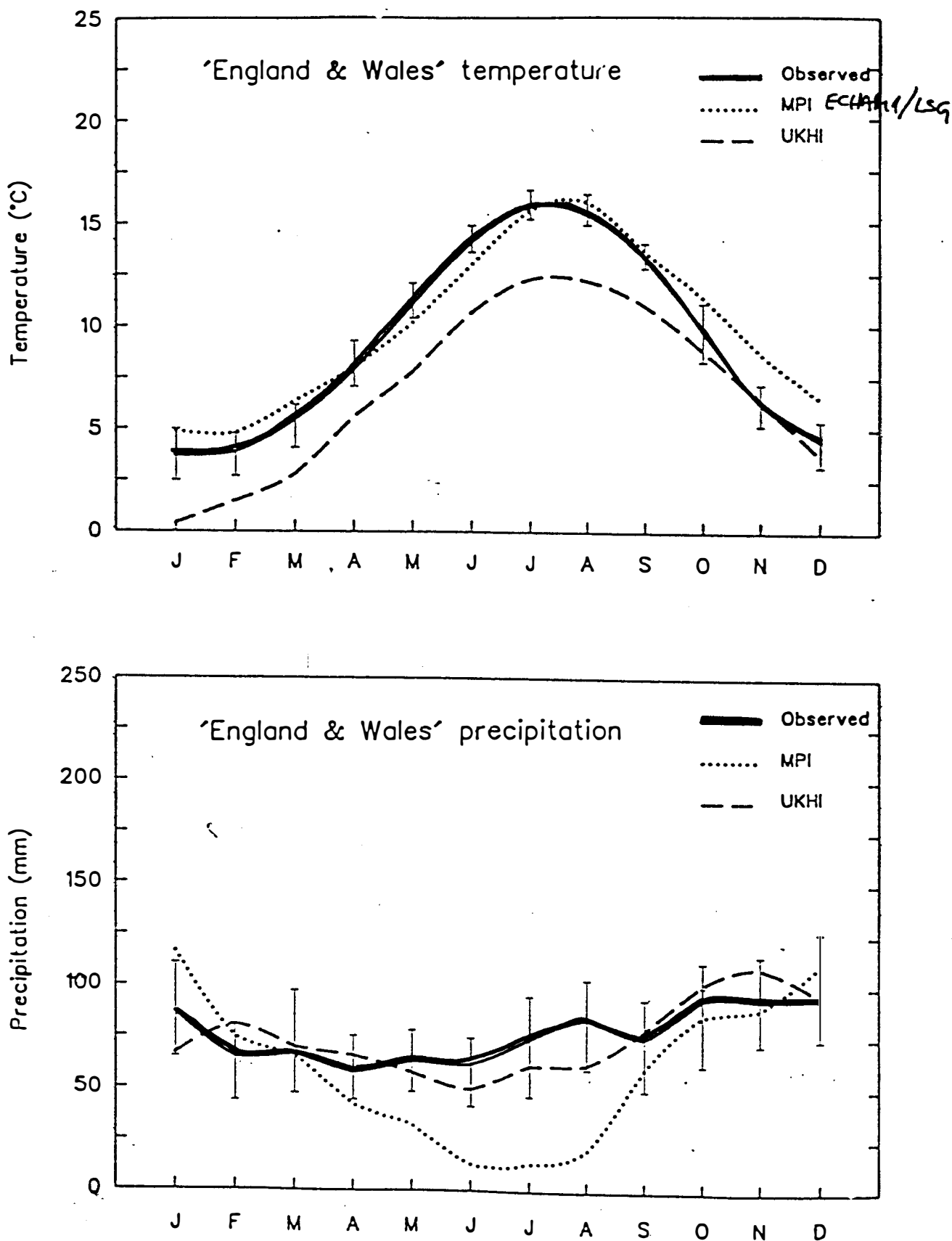


Figure 5: Observed monthly mean temperature and precipitation over an area representative of England & Wales and simulated data for a 5° gridbox centred at 52.5°N and 2.5°W (models; see Figure 1). The observed temperature data are from Manley (1974; updated by Parker *et al.*, 1992) and the precipitation data from Wigley *et al.* (1984; updated by Gregory *et al.*, 1991). The range bars on the observed lines are calculated as in Figure 3 from the 110-year observed records from 1881 to 1990.

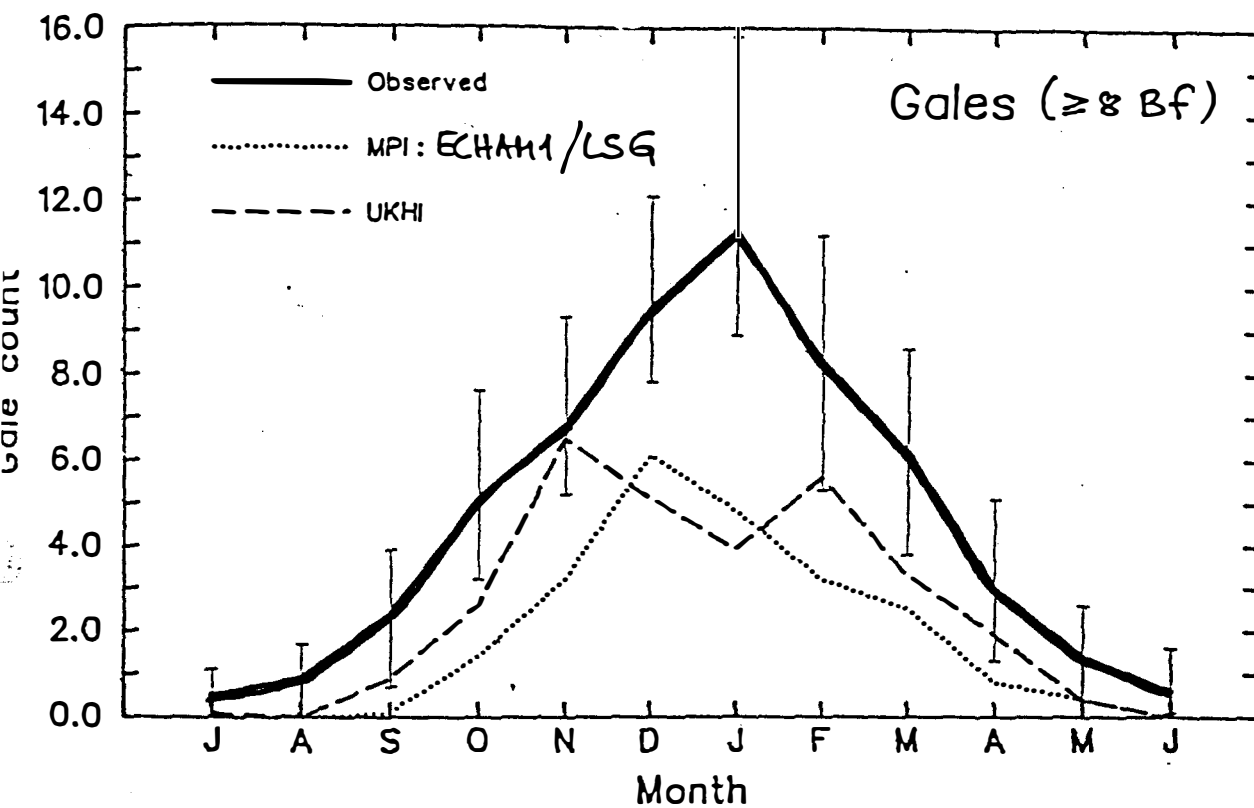


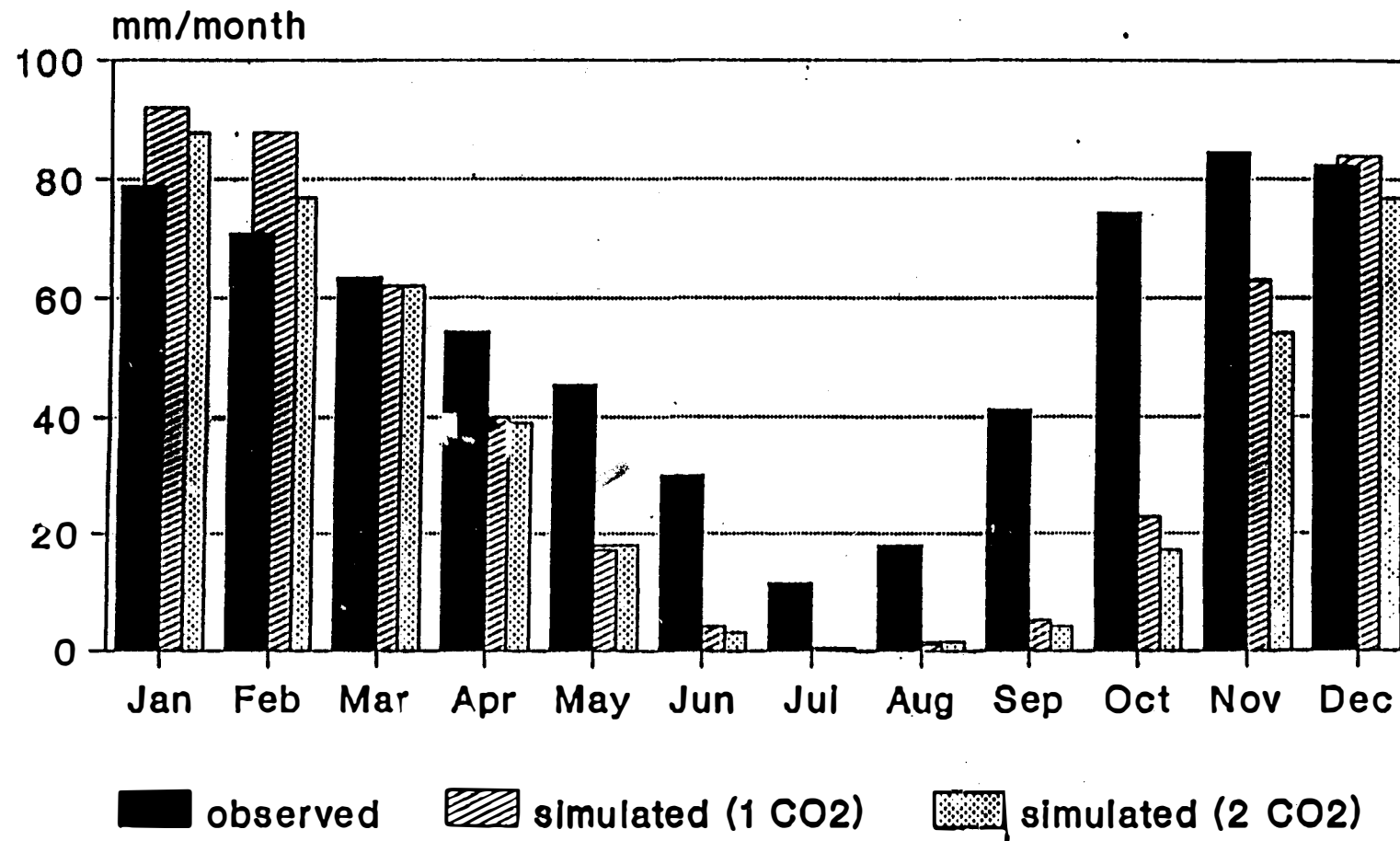
Figure 6: Observed and model-simulated mean monthly frequencies of Gales over the British Isles. The range bars on the observed line are calculated as in Figure 3 from the 110-year observed record from 1881 to 1990. Note that the 'gale year' runs from July to June.

*Hulme et al, 1992*



# Iberian Rainfall

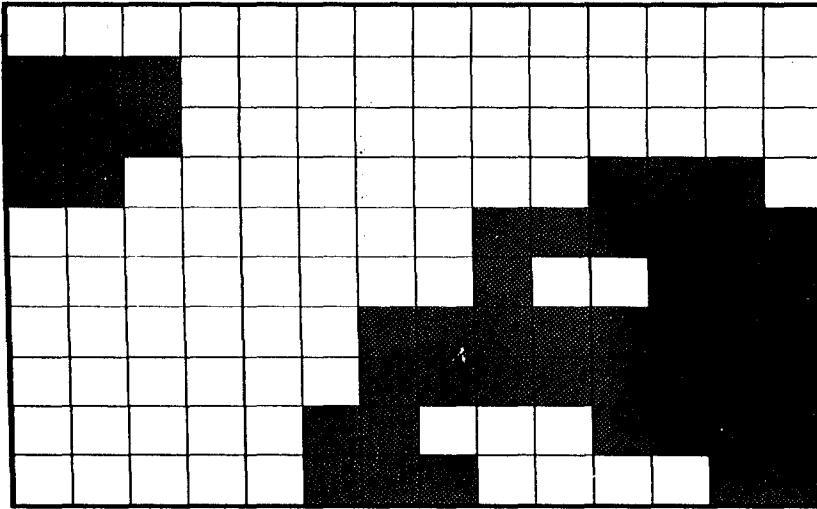
## Observed and Simulated



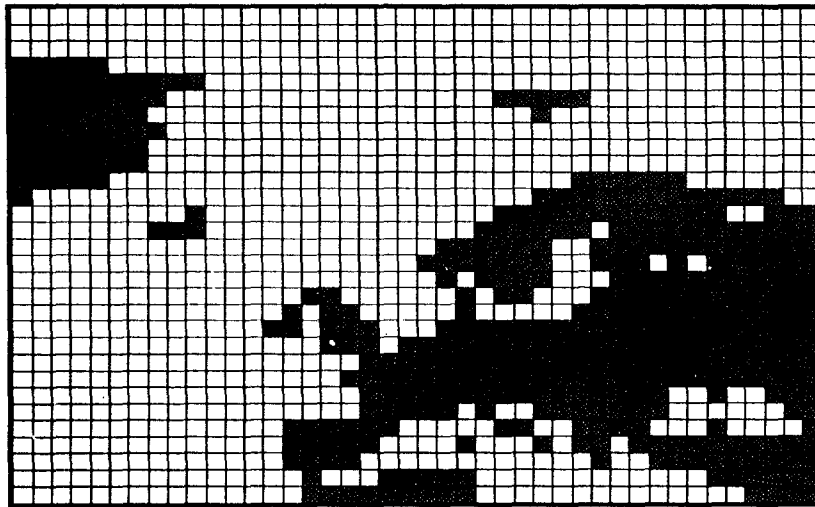
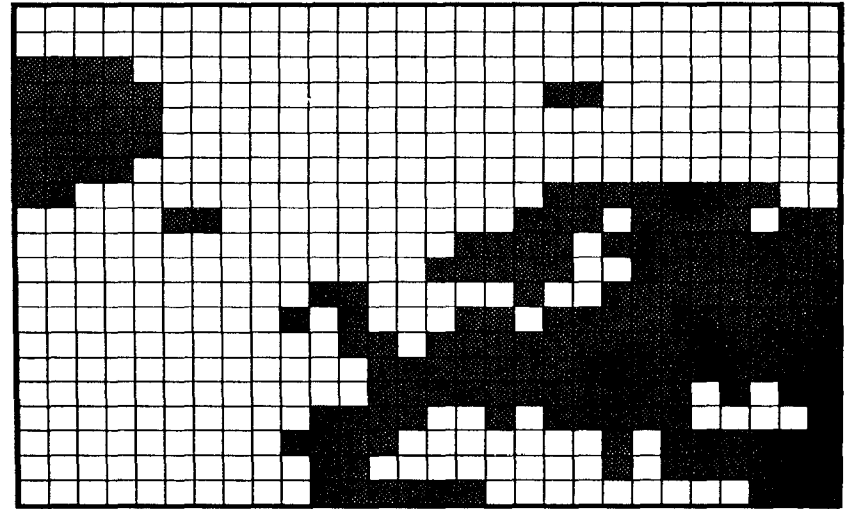
ECHAM

# European part of the land-sea mask for different T-model resolutions

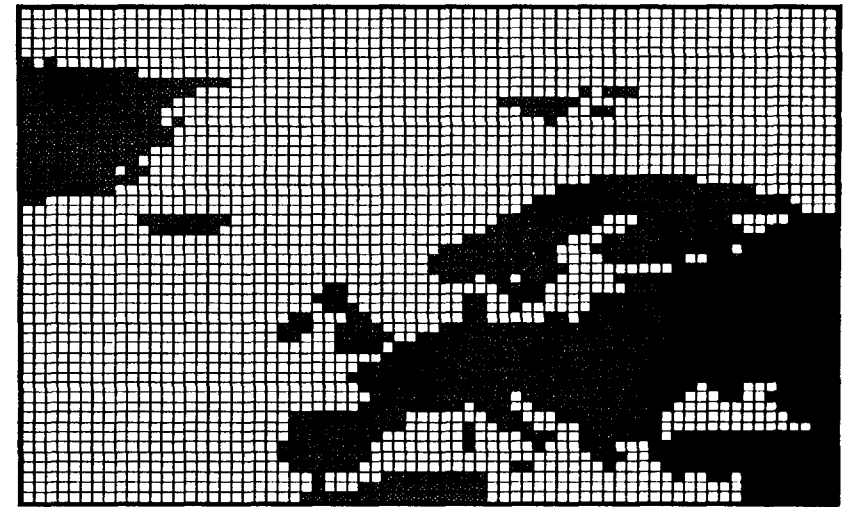
a) T21



b) T42



c) T63

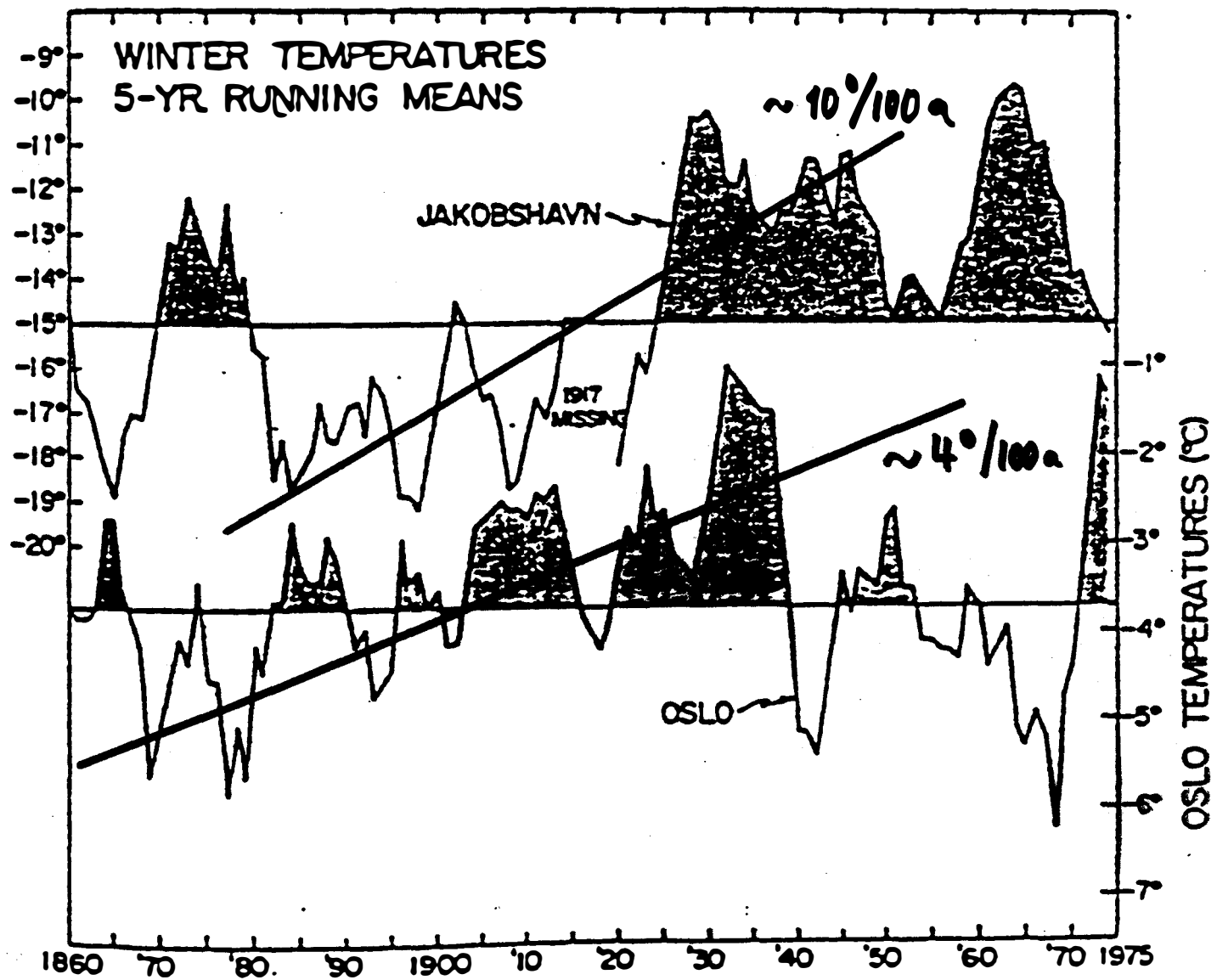


d) T106

TOP



from: Chris Huntington 1992

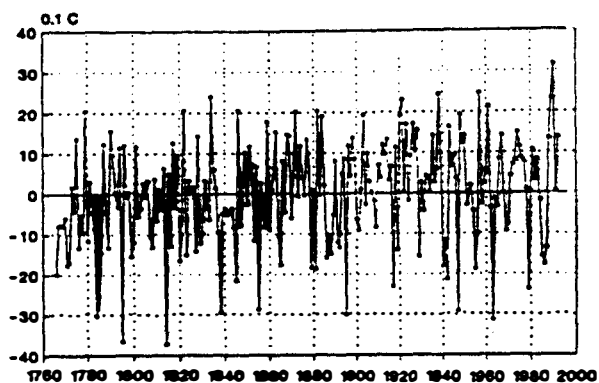


Five-year running means of winter temperature at Jakobsbavn and Oslo ( $^{\circ}C$ ).

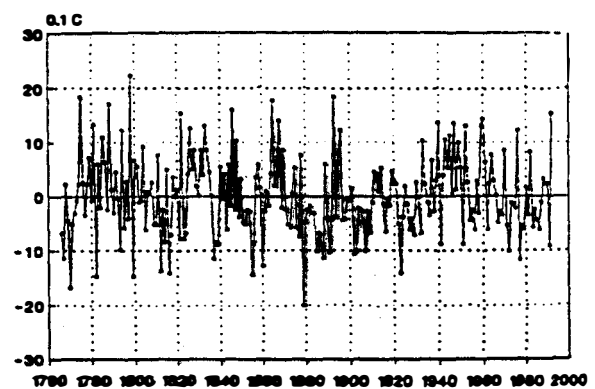
Case 2 : Significant warming trend  
in N Europe ... no public response  
(Arrhenius'  $CO_2$ -theory  $\sim 1898$ )

van Loon & Rogers  
1978

Mean of January-February-March

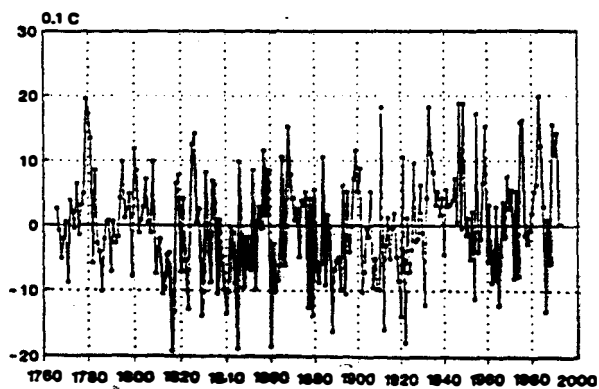


Mean of April-May-June

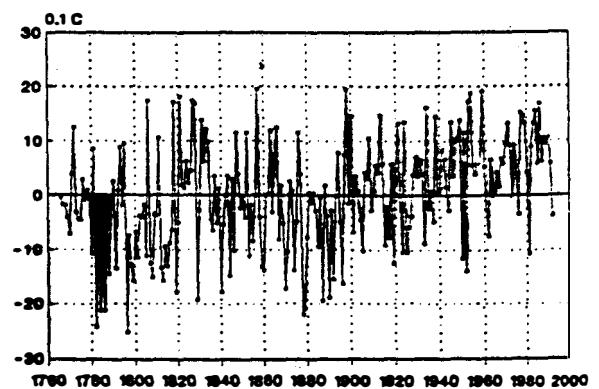


## Middle England Temperature

Mean of July-August-September



Mean of October-November-December



files: desktop, \hg\book\eng-XXX, XXX=jfm,amj,jas,ond hvs

# The Great Challenge

Climate impact research is an interdisciplinary business. The fundamental problem is the transfer of information across the borderline of otherwise well separated (and even well entrenched) “science states”.

- **Today**, and to some extent for the past, we can have information with high space-time accuracy, so that the climate-related input into (mathematical or conceptual) models is no problem.
- **Scenarios for the future**, however, are inherently uncertain with respect to quantities and space and time. This uncertainty is not the result of unable climate researchers but a fundamental problem.
- **Climate impact models** (independently whether they are mathematical models run on a computer or conceptional models) must meet this uncertainty of the input. Models useful for the representation of the present state are not necessarily equally useful for the assessment of future changes. Possibly, models have to be de-complexified in order to serve as an climate impact assessment tool.