



Introduction & Motivation

- □ Storms are a characteristic of the maritime weather at midlatitudes. It is particularly difficult to homogeneously "measure" storm intensities:
- Observations depend on subjective decisions
- Long series of direct wind observations often suffer from inhomogeneities
- □ It seems that Northern European storm climate has been relatively stable during the last two hundred years (Bärring and von Storch 2004). However, in history many people have experienced storms as having become more severe (von Storch and Stehr 2000).
- □ From proxy data we know that past climate has undergone significant variations: \rightarrow Medieval Warm Period, Little Ice Age, Late Maunder Minimum (LMM, 1675-1710)
- □ The spatial extent of these episodes is still a subject for debate. But there is little doubt that significant temperature deviations prevailed during these times in large areas.

Are these temperature variations associated with significant changes of storminess?

Climate Model

Coupled Global Ocean-Atmosphere General Circulation Model ECHAM4/HOPE-G Atmosphere: horizontal resolution 3.75°x3.75° (T30), vertical: 19 layers. Ocean: horizontal resolution 2.8°x2.8°(T43), refining to 0.5° near equator, vertical: 20 layers.

Flux correction is applied.

Simulations

- 2 realizations of a historic climate change simulation (time dependent solar, volcanic, and anthropogenic GHG forcing). **One simulation is shown exemplarily.**
- 3 transient climate change simulations with idealized forcing simulating a future climate state (CMIP2, SRES A2/B2 scenarios). The A2/B2 scenarios are shown exemplarily.
- a control simulation (present day conditions) (Zorita et al, 2004)



References

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Modelling the variability of midlatitude storm activity on decadal to century time scales

(e.g. Roeckner et al. 1996)

Effective solar forcing and annual global mean temperature. The temperature shows two distinct minima: one during the LMM and one during the Dalton Minimum.

- Methods Historical times and industrially influenced period: □ Variations of the external forcing are not directly related to variations in storm activity, neither in historical times, nor in the industrially influenced period. □ Characteristic parameter: Frequency of **maximum wind speed** □ In pre-industrial times, storm activity is decoupled from hemispheric mean temperature. events. Lower limit on the Beaufort wind speed scale of 8 Bft_ Anomalous temperature regimes (like the LMM) are not associated with systematic is used as threshold (17.2 ms⁻¹). storm conditions. Determination of 2 simple indicators describing the frequency □ The storm indices show no trend until recently (except for SH in 20th century). and regional shift of storm activity: Future climate change scenario: □ Poleward shift of regions with maximum storm intensity for N-Atl, N-Pac and SH. Storm frequency index: Simple count of storm days per grid point □ Storm intensity is constant over the NH as a whole, but increasing in the N-Atl and • Storm shift index: PC obtained by projection of storm frequency decreasing in the N-Pac region. data onto dominating EOF of storm frequency in pre-industrial Parallel increase of storm indices and temperatures over the N-Atl region and times Southern Ocean. (Fischer-Bruns and von Storch, 2005) Results

- Analysis of the simulations with respect to global storm activity



Storm frequencies for the industrially influenced period hardly differ from pre-industrial maps, even though significant temperature anomalies temporarily emerge in the historical



and differences between future period and pre-industrial period



Storm frequency and future shift

Frequency of storm days for DJF and JJA for the pre-industrial period 1551–1850 in one of the historical simulations

In the climate change simulations, a poleward shift of storm activity is found in all three storm track regions. Here results are shown for the IPCC scenario A2:





lines.





Conclusions

